

Dust Options in WRF-Chem

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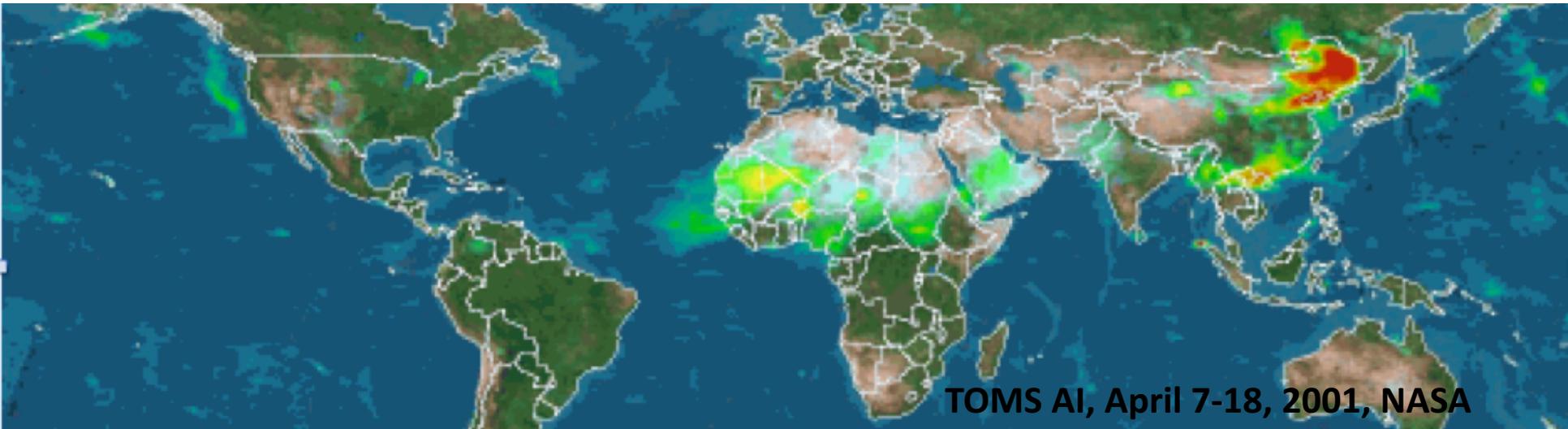
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Boulder, CO, USA*

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Several Dust Options in WRF-Chem

Challenges in Estimating Dust Emissions

Understanding changes in dust emissions is important for both interpreting past and predicting future climate change.



- **Large uncertainties in estimating global dust emissions** that the models simulate global dust emissions between 514 and 4313 Tg yr⁻¹ and dust loads ranging from 6.8 to 29.5 Tg (Textor et al., 2006; Huneus et al., 2011).
- The magnitude of dust emissions to the atmosphere depends on the surface wind speed and soil characteristics, thus the spatial and temporal variability **is easily influenced by changes in regional and global meteorological fields and surface properties.**
- Incompletely understanding of the physical processes that determine the emitted dust in model, **e.g., threshold fraction velocity, horizontal saltation flux and vertical flux**

Current WRF-Chem Dust Options

1. GOCART dust scheme (**dust_opt= 1**)

module_gocart_dust.F

2. AFWA dust scheme (**dust_opt=3**)

module_gocart_dust_afwa.F

3. UoC dust scheme (**dust_opt=4**)

dust_schme=1 (Shao 2001)

dust_schme=2 (Shao 2004, S04)

dust_schme=3 (Shao 2011, S11)

dustwd_onoff=0 (turn off Jung 2004 dust wet deposition)

dustwd_onoff=1 (turn on Jung 2004 dust wet deposition)

module_uoc_dust.F

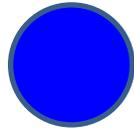
module_uoc_dustwd.F

module_qf03.F

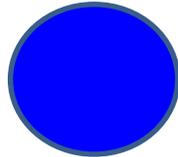
Dust Size Bins in WRF-Chem

radii

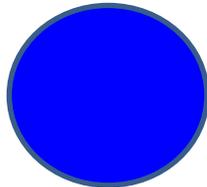
0– 1.0 μm
r-eff=0.73



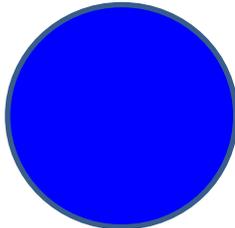
1.0–1.8 μm
r-eff=1.4



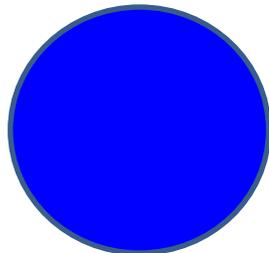
1.8–3.0 μm
r-eff=2.4



3.0–6.0 μm
r-eff=4.5



6.0–10.0 μm
r-eff=8.0



Dust in WRF-Chem is distributed into 5 size bins:

Type	reff_dust(μm)	Density(g cm ⁻³)
bin1	0.73	2500
bin2	1.4	2650
bin3	2.4	2650
bin4	4.5	2650
bin5	8	2650

- $\text{PM}_{2.5} = \text{bin1} + 0.3125 * \text{bin2}$
- $\text{PM}_{10} = \text{bin1} + \text{bin2} + \text{bin3} + 0.87 * \text{bin4}$

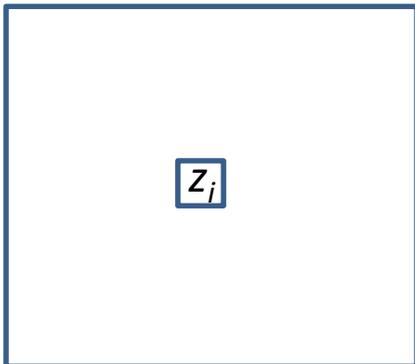
GOCART Dust Source Function S

- Source function S is the fraction of alluvium available for wind erosion, as follows:

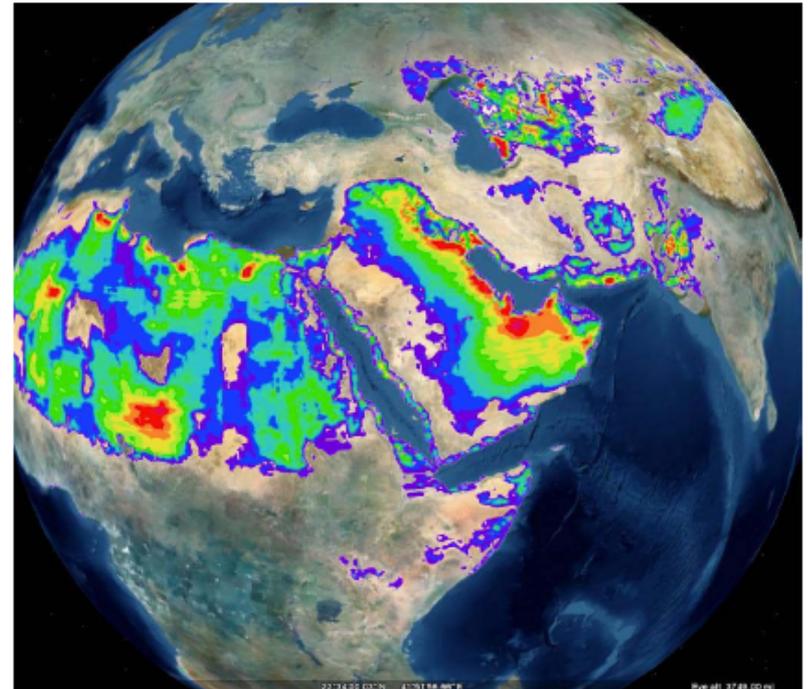
$$S = \left(\frac{z_{\max} - z_i}{z_{\max} - z_{\min}} \right)^5 ,$$

S : the probability to have accumulated sediments in the grid cell i of altitude z_i ,

z_{\max} and z_{\min} : the maximum and minimum elevations in the surrounding 10x10 degree topography, respectively.



Ginoux Erodibility Factor



GOCART Dust Scheme

Dust uplifting flux (equivalent empirical formulation by Gillette and Passi [1998]):

$$F_p = C_G S s_p U_{10}^2 (U_{10} - U_t^*), \quad U_{10} > U_t^*$$

Dimensional tuning constant horizontal wind speed at 10m

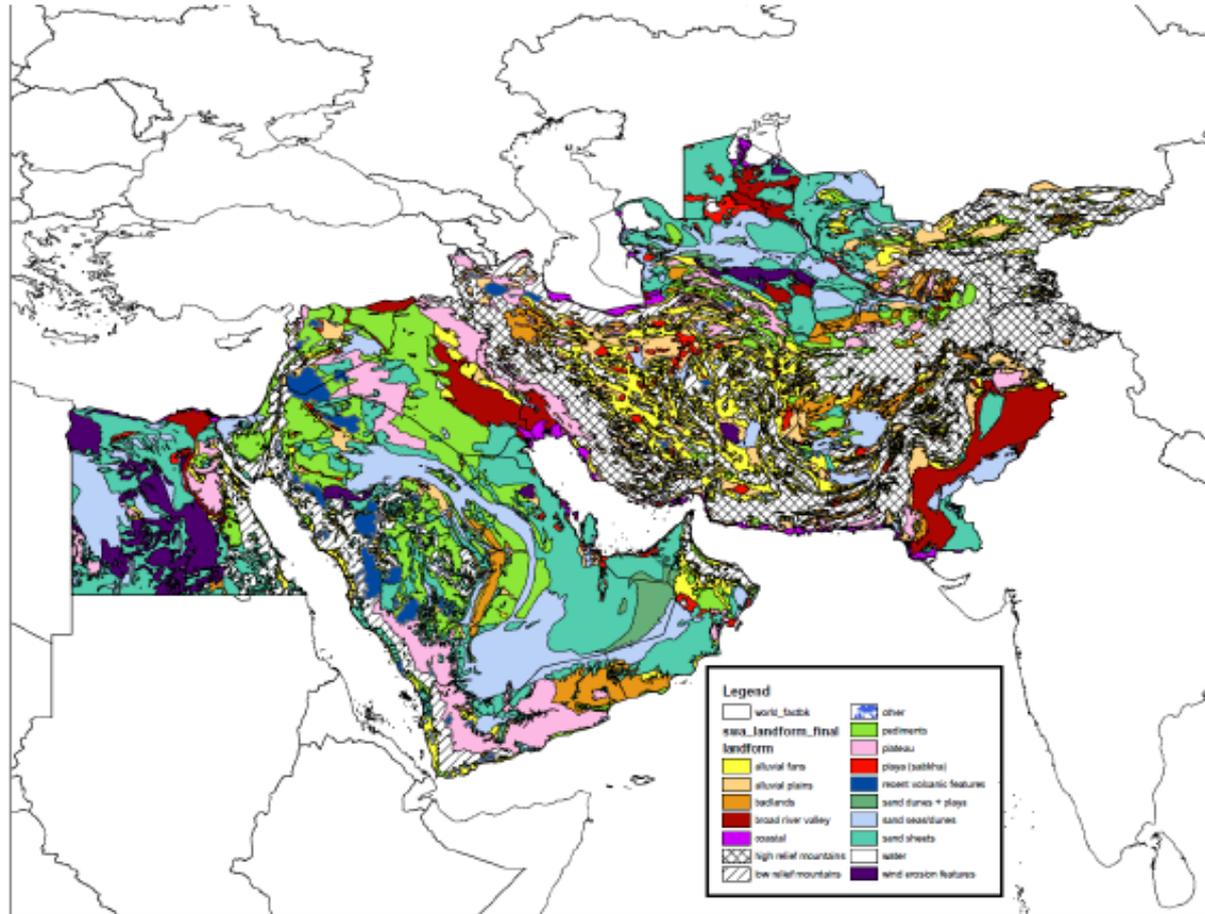
Erodibility function fraction of each size class threshold velocity

- Ginoux et al., 2001: The s_p values are thus 0.1 for the class 0.1-1 μm , and 1/3 for the classes 1-1.8 μm , 1.8-3 μm , and 3-6 μm , respectively.
- Ginoux et al., 2004: s_p values are 0.1, 0.25, 0.25, 0.25, 0.25

- ✧ The threshold velocity (U_t^*) is actually calculate the threshold friction velocity, but they are not the same physical meaning even the same unit.
- ✧ U_t^* is original from Bagnold (1941) and now from Marticorena and Bergametti (1995) in WRF-Chem.

Erodibility in AFWA

Southwest Asia Landform Map (Desert Research Institute (DRI))



AFWA and DRI developing physical process based erodibility database!

AFWA Dust Scheme

Saltation Flux Over Bare Soil (Kawamura, 1951):

$$H(D_p) = C \frac{\rho_a}{g} u_*^3 \left(1 + \frac{u_{*t}}{u_*}\right) \left(1 - \frac{u_{*t}^2}{u_*^2}\right) \quad G = \sum H(D_p) dS_{rel}(D_p)$$

Using friction velocity u_ instead of using horizontal wind speed at 10m as GOCART scheme, that be consist with the threshold friction velocity u_{*t} .*

Bulk vertical dust flux scheme that based on Marticorena and Bergametti (1995):

Bulk Vertical Dust Flux (efficiency factor (α): Gillette, 1979)

$$F_{bulk} = G\alpha \times \text{Erod}$$

$$\alpha = 10^{0.134(\% \text{clay}) - 6}$$

■ **Threshold Friction Velocity (Iversen & White, 1982)):**

$$u_{*t}(D_p) = 0.129 \frac{\left[\frac{\rho_p g D_p}{\rho_a}\right]^{0.5} \left[1 + \frac{0.006}{\rho_p g D_p^{2.5}}\right]^{0.5}}{\left[1.928(aD_p^x + b)^{0.092} - 1\right]^{0.5}}$$

$$u_{*t} = u_{*t}(D_p) \frac{f(\text{moisture})}{f(\text{roughness})}$$

AFWA Dust Scheme

- Correction factors applied to u_{*t} ,

$$u_{*t} = u_{*t}(D_p) \frac{f(\text{moisture})}{f(\text{roughness})}$$

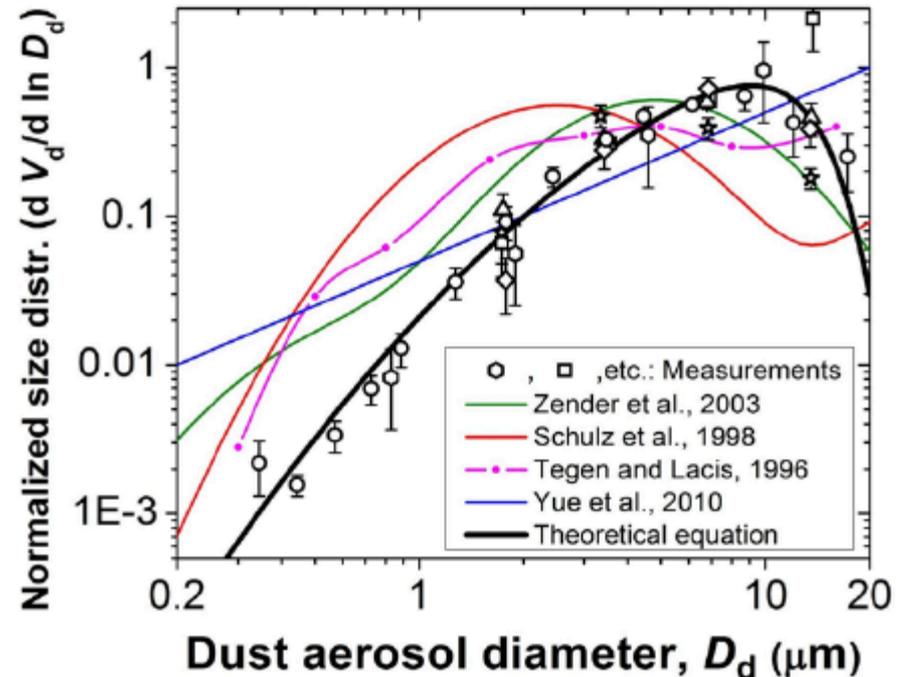
- $f(\text{roughness})$ is a drag partition correction
 - $f(\text{roughness}) = 1.0$ implies the surface is smooth; value decreases with increasing amounts of large rocks, cobbles, vegetation, etc.
 - Currently set to 1.0; representative of Southwest Asia.
 - Dust emission is restricted to areas with roughness length $z_0 \leq 5\text{m}$ (typically barren lands and sparsely vegetated surfaces).

Dust Size Distribution in AFWA Scheme

- Marticorena and Bergametti scheme only provides bulk dust flux.
- Particle Size Distribution (PSD) developed by Dr. Jasper Kok (NCAR)
 - Brittle material fragmentation theory
- Kok, 2010 (PNAS)

Fraction of Five Size distributions

Type	GOCART	AFWA
bin1	0.1	0.1074
bin2	0.25	0.1012
bin3	0.25	0.2078
bin4	0.25	0.4817
bin5	0.25	0.1019



Dust Size Redistribution

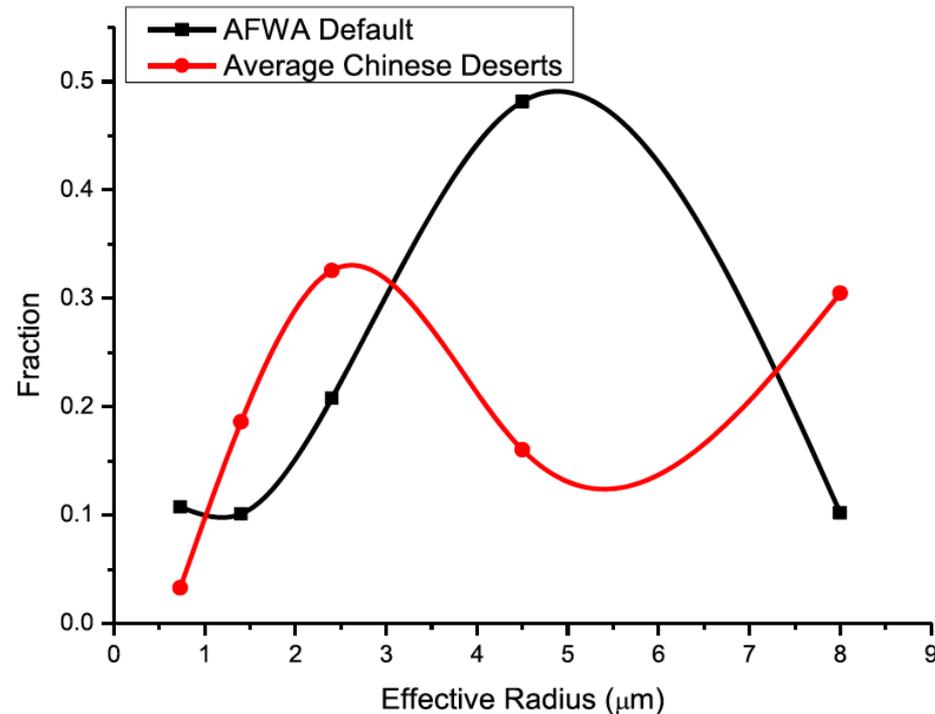
bin1+...+bin5



Total emissions of all bins



Redistribute the emission of each size bins based on the new particle size distribution fraction.



Su and Fung 2015

For AFWA,
phys/module_data_gocart_dust.F,
the variable of size distribution
fraction is 'distr_dust', that can be
modified to your own fraction from
the operation

UOC Dust Scheme (dust_schme=2 or dust_schme=3)

Shao [2004] proposed a new size-resolved dust emission scheme (S04):

$$F(d_i, d_s) = c_y \eta_f [(1 - \gamma) + \gamma \sigma_p] (1 + \sigma_m) g \frac{Q_{ds}}{u_*^2},$$

- $F(d_i, d_s)$: dust emission rate of size d_i generated by saltation of size d_s
- C_y : dimensionless coefficient; η_f : fraction of emitted dust
- σ_m : bombardment efficiency
- σ_p : ratio between fraction of free dust and fraction of aggregated dust
- γ : function that describes how easily aggregated dust can be released
- Q_{ds} : saltation flux of size d_s
- g : acceleration due to gravity

The erodibility factor is only used to constrain the potential emission regions instead of being used to scaling the dust emission directly as in AFWA scheme and GOCART scheme.

S04 (Shao 2004) and S11 (2011)

The emission of dust of size d_i associated with the saltation of all grain sizes can be estimated as a weighted average over the sand particle size range defined by d_1 and d_2 :

$$F(d_i) = \int_{d_1}^{d_2} F(d_i; d) p_s(d) \delta d$$

Total dust emission:

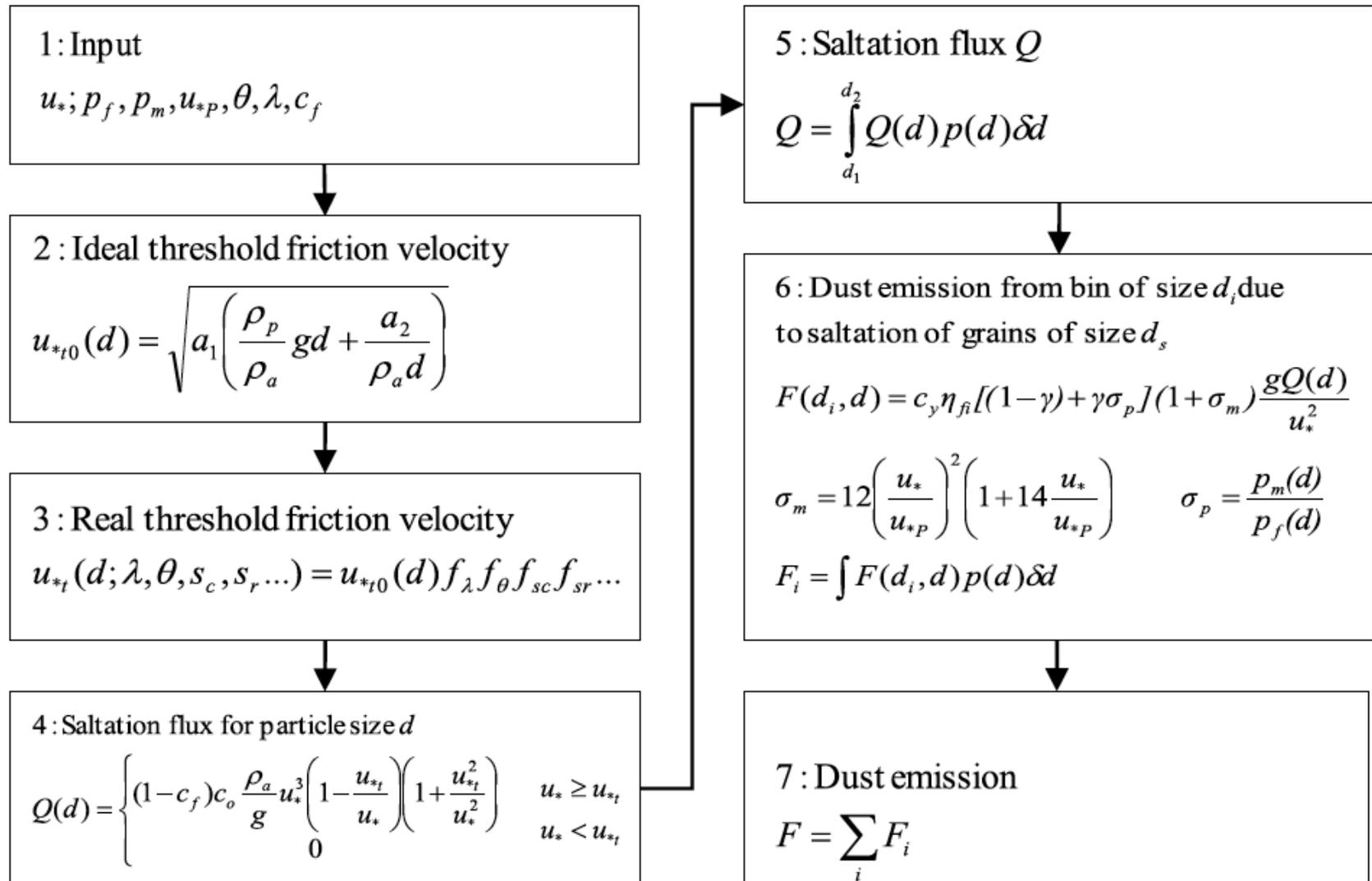
$$F = \sum_{i=1}^I F(d_i)$$

This scheme makes use of $p_m(d)$, the minimally disturbed psd, and $p_f(d)$, the fully disturbed psd, of the parent soil to constrain the size distribution of the airborne sand and dust particles,

$$p_s(d) = \gamma p_m(d) + (1 - \gamma) p_f(d)$$

The S11 scheme is a simplification of S04 with $\gamma = 1$, which means that $p_f(d)$ is no longer necessary in the simplified scheme.

Structure of the Wind erosion Scheme (S04 and S11)



- Simulation of a dust storm in Australia with Shao (2004) dust emission scheme (dust_opt = 4, dust_schme = 2).

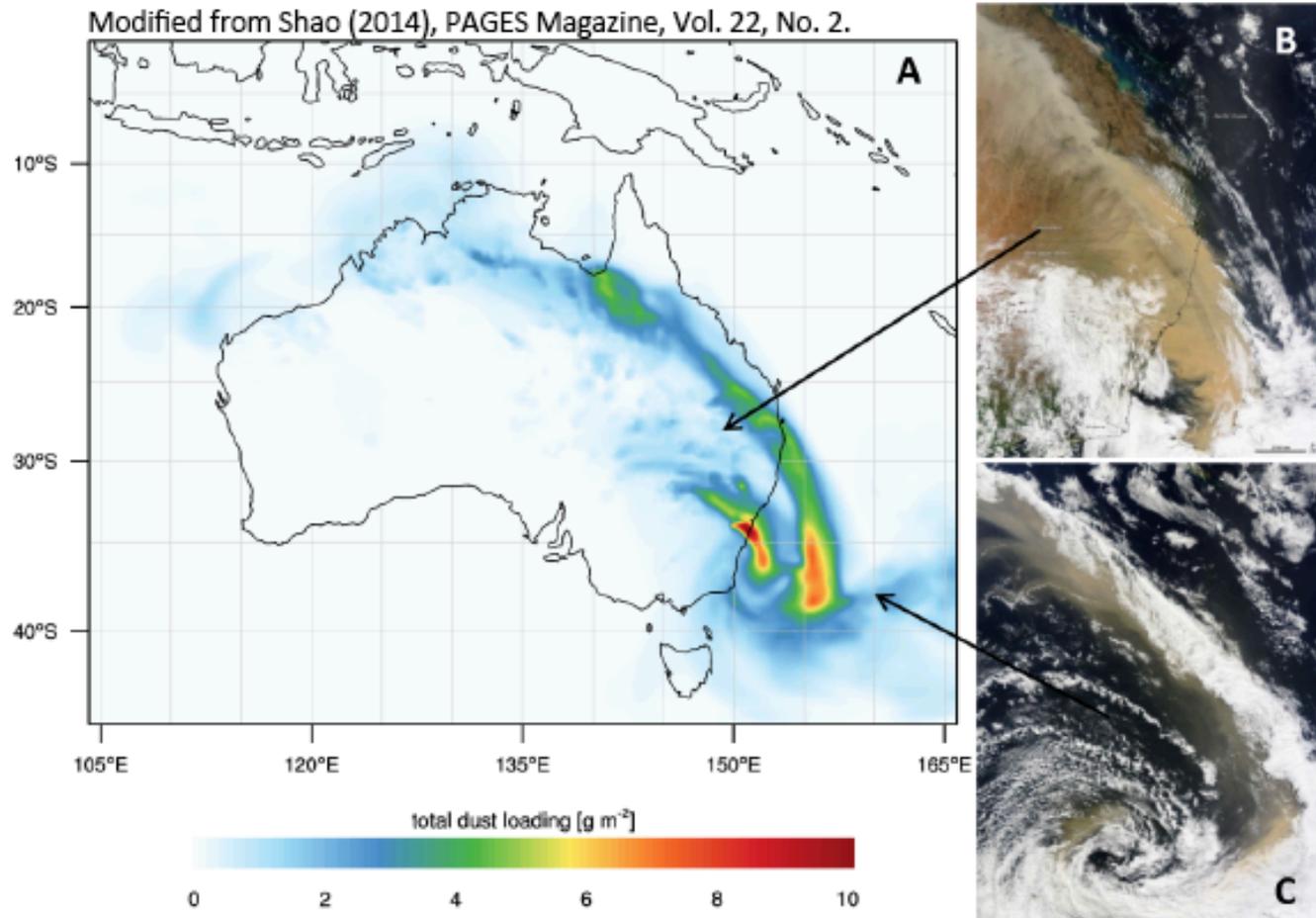


Fig. 1: (A) Dust load on 23 September 2009 modeled with WRF-Chem at 30km horizontal resolution; MODIS satellite images from (B) 23 September 2009 and (C) 24 September 2009.

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Major Dust Scheme References in WRF-Chem

dust_opt = 1 (GOCART dust emissions):

- Ginoux, P., M. Chin, I. Tegen, J. M. Prospero, B. Holben, O. Dubovik, and S.-J. Lin , 2001: Sources and distributions of dust aerosols simulated with the GOCART model. *J. Geophys. Res.*, **106(D17)**, 20255-20273.

dust_opt = 3 (GOCART with AFWA modifications):

- Jones, S. L, Adams-Selin, R., Hunt, E. D., Creighton, G. A., Cetola, J. D., 2012: Update on modifications to WRF-CHEM GOCART for fine-scale dust forecasting at AFWA. *AGU Fall Meeting Abstracts*.
- Jones, S. L, Adams-Selin, R., Hunt, E. D., Creighton, G. A., Cetola, J. D., 2010: Adapting WRF-CHEM GOCART for Fine-Scale Dust Forecasting. *AGU Fall Meeting Abstracts*, Vol. 1.

dust_opt = 4 (GOCART with UoC modifications):

- Shao, Y, 2001: A model for mineral dust emission. *J. Geophys. Res.*, **106**,20,239-20,254.
- Shao, Y, 2004: Simplification of a dust emission scheme and comparison with data *J. Geophys. Res.*, **109**, doi:10.1029/2003JD004372.
- Shao, Y., M. Ishizuka, M. Mikami, J. Leys , 2011: Parameterization of size-resolved dust emission and validation with measurements. *J. Geophys. Res. Atmos.*, **116**, D08203, doi:10.1029/2010JD014527.

Some notes

- ❑ Be careful with the dust emission unit in the output, make sure it has been multiply the grid area when calculate global total.
- ❑ When getting initial or boundary condition from other model, make sure that the dust particle size are consistent, otherwise, better to get the close fraction of dust concentrations for each size bins. The AOD calculation is very sensitive to the fraction of dust concentrations of each size bins, when compare with other model, be careful with the differences within it.
- ❖ Other questions, getting information from WRF-Chem web site: <https://ruc.noaa.gov/wrf/wrf-chem/> or email group: <https://ruc.noaa.gov/wrf/wrf-chem/forum.htm>
- ❖ Free to contact us via wrfchemhelp.gsd@noaa.gov